

Does water escape from subducted slabs at the core-mantle boundary?

Hydrated SiO_2 phase is the main carrier of water in the subducting slabs of the Earth's lower mantle. Recent experiments performed along cold-to-normal lower-mantle geotherms have demonstrated that SiO_2 can hold up to ~ 3.5 wt% H_2O [1,2]. Regardless of the high water holding capacity of dense SiO_2 , it has been speculated that subducted slabs will dehydrate at high temperatures in the core-mantle boundary (CMB) region, and the released water will give rise to chemical heterogeneities, as indicated by seismological observations. Water can induce melting, melt migration, and hydration of the lowermost mantle materials, creating large low-shear-velocity provinces. Furthermore, the reaction of water with the core metal can form superoxidized FeOOH_x ($x < 1$), which accounts for ultralow velocity zones. Nevertheless, the solubility and partitioning of water into Al-bearing hydrous SiO_2 within subducted former crustal rocks, especially mid-oceanic ridge basalt (MORB), have not been investigated under the high pressure and temperature (P – T) conditions in the CMB region; therefore, the dehydration of subducted slabs and its associated effects remain unverified.

We conducted 19 separate melting experiments on hydrous MORB up to the lowermost mantle pressure

range at 4100 K in a laser-heated diamond-anvil cell by collecting X-ray diffraction (XRD) patterns at SPing-8 BL10XU [3]. Melting textures and chemical compositions, including the H_2O content in the melt and coexisting solids, were examined in the recovered samples. They exhibited a concentric texture: a round pocket of quenched partial melt at the center surrounded by a SiO_2 phase, CaSiO_3 perovskite (davemaoite), SiO_2 – AlOOH solid solution (ss.) and bridgmanite at 25–59 GPa (Fig. 1(a)). At 100 GPa and above, SiO_2 was still present next to the melt pool, and the bridgmanite/post-perovskite appeared closer to the melt than it did at lower pressures (Fig. 1(b)). The XRD patterns of these samples indicated that CaCl_2 -type SiO_2 appeared down to at least 25 GPa at 2900 K. This observation is consistent with earlier findings that CaCl_2 -type SiO_2 stabilizes over stishovite in the presence of Al_2O_3 and H_2O [2], whereas it forms above 65–75 GPa at 1500–2000 K in Al-free dry SiO_2 and SiO_2 – H_2O systems [1]. Furthermore, the crystal structure of SiO_2 transformed from CaCl_2 -type to α - PbO_2 -type (seifertite) above 128 GPa in the lowermost mantle (Fig. 2).

Water concentrations in the SiO_2 phase, SiO_2 – AlOOH ss., and partial melts were determined by

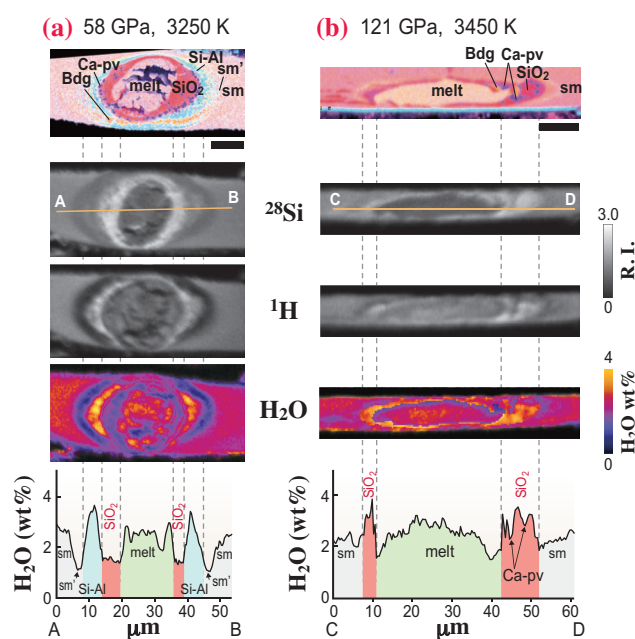


Fig. 1. Cross sections of partially molten samples. Samples were recovered from (a) 58 GPa and 3250 K and (b) 121 GPa and 3450 K. (Uppermost panels) EDS X-ray elemental maps combined for Si, Al, Ca, and Mg, showing melt, hydrated SiO_2 (red), CaSiO_3 perovskite (Ca-pv; blue), SiO_2 – AlOOH ss. (Si-Al; light blue), bridgmanite (Bdg; orange), and dehydrated (sm') and original starting material (sm). (Middle panels) Corresponding secondary ion images for $^{28}\text{Si}^+$ and $^1\text{H}^+$ and the distribution maps of H_2O . Scale bar: 10 μm .

secondary ion mass spectrometry coupled with high-resolution imaging (Fig. 1). The water content in Al-bearing CaCl_2 -type SiO_2 substantially increased with increasing pressure, reaching 3.6 wt% H_2O at 121 GPa under 3450 K (Figs. 2(a,b)). This value is considerably higher than the recent H_2O solubility estimate of only 0.4 wt% under the same P – T condition based on experiments on the Al-free SiO_2 – H_2O system [1]. This discrepancy suggests that the charge-coupled substitution $\text{Si}^{4+} = \text{Al}^{3+} + \text{H}^+$ is the key mechanism for water incorporation into Al-bearing CaCl_2 - [2] and α - PbO_2 -type SiO_2 (Fig. 2(c)). Our XRD study found that the magnitude of the orthorhombic distortion increased with increasing pressure (Fig. 2(d) and Fig. 3), and this crystallographic distortion correlated strongly with the water content (Fig. 2(e) and Fig. 3). The enhanced distortion shortened the distance between specific pairs

of oxygen atoms, making the hydrogen bonds stronger and stabilizing hydrogen in SiO_2 . Moreover, the overall increase in water concentration in CaCl_2 -type SiO_2 with increasing pressure resulted in an increase in the partition coefficient of water ($D_{\text{H}_2\text{O}}$ (CaCl_2 – SiO_2 /melt)) from ~ 0.2 at 25 GPa to 1.4 under deep lower-mantle conditions at 120 GPa (Fig. 2(f)). The H_2O content in Al-bearing α - PbO_2 -type SiO_2 , formed at 3650–4100 K and 128–144 GPa—conditions corresponding to the Earth's CMB region—remained nearly constant at 2 wt% (Fig. 2(a)). Moreover, $D_{\text{H}_2\text{O}}$ (α - PbO_2 – SiO_2 /melt) ranged from 1.1 to 2.1 (Fig. 2(f)), indicating that water preferentially partitions into hydrated SiO_2 rather than into coexisting partial melts under CMB conditions.

The dehydration of subducted slabs at the base of the mantle has been highlighted, and its consequences have been extensively discussed. However, our discovery of the high H_2O storage capacity—approximately 2 wt% in α - PbO_2 -type SiO_2 and ~ 0.3 – 0.6 wt% in subducted MORB crust—suggests that, in practice, water does not escape from slabs even under the high temperature conditions of the CMB region.

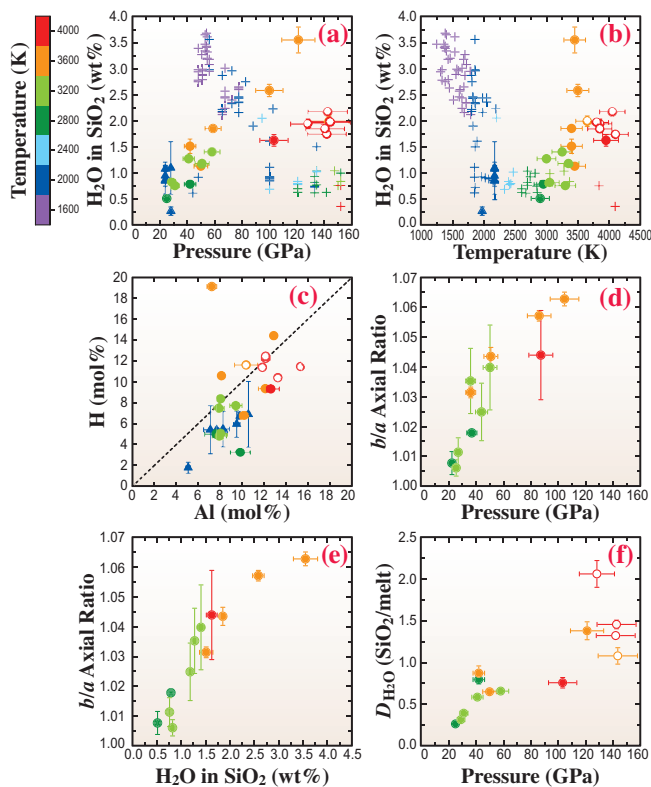


Fig. 2. Variations in hydrated SiO_2 . The present experiments on hydrous MORB show data for CaCl_2 -type (closed circles) and α - PbO_2 -type SiO_2 (open circles). Earlier data obtained in water-saturated SiO_2 – H_2O (pluses) [1] and SiO_2 – Al_2O_3 – H_2O systems (triangles) [2] are also given. Color indicates temperature. Water concentrations in SiO_2 , plotted as functions of (a) pressure and (b) temperature. (c) Correlation between Al and H concentrations in SiO_2 , indicating that the charge-coupled substitution $\text{Si}^{4+} = \text{Al}^{3+} + \text{H}^+$ is a main mechanism for water incorporation. (d) b/a axial ratio of the CaCl_2 -type structure observed at 300 K, representing the magnitude of the orthorhombic distortion from a tetragonal structure with $b/a = 1$. (e) Correlation between the H_2O content and b/a ratio. (f) SiO_2 /melt partition coefficient of H_2O , demonstrating that water is preferentially partitioned into SiO_2 rather than into silicate melt in the lowermost mantle.

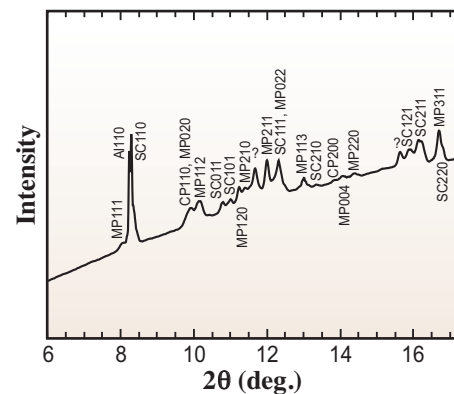


Fig. 3. XRD patterns of the solid part around a melt pocket obtained after quenching. SC, hydrated CaCl_2 -type SiO_2 ; MP, bridgmanite; CP, CaSiO_3 perovskite; Al, SiO_2 – AlOOH ss. Peak separations of SC121/211 and SC011/101 indicate the orthorhombic distortion of the CaCl_2 -type structure.

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